

土壤呼吸作用时空动态变化及其影响机制研究与展望

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摘要 测定不同陆地生态系统土壤呼吸速率及其时空波动, 阐明其影响因子, 对于全球碳素平衡预算和全球变化潜在效应估计是最为基本的数据。然而, 有关土壤呼吸作用变异性及其影响因素的知识仍存在局限性, 一些关键的过程和机制还有待阐明。该文综述了近年来土壤呼吸作用时空动态规律、影响机制和模拟方面的研究进展, 指出环境因子和生物因子共同驱动着土壤呼吸作用的时间动态变化; 土壤呼吸作用在不同时间尺度上还具有明显的空间异质性, 这主要是植被覆盖、根系分布、主要的环境因素和土壤特性空间分布的异质性造成的。生物因子是影响土壤呼吸作用时空动态变化的主要因素之一。然而, 目前所使用的土壤呼吸作用经验模型通常利用土壤温度、土壤湿度或者两者的交互作用模拟土壤呼吸作用动态变化, 但没有考虑生物因子的影响, 这可能会导致明显的偏差和错误。因此, 为了精确估算土壤呼吸作用, 必须解决土壤呼吸作用小尺度上的空间变异性; 加强不同时间尺度上生物要素对土壤呼吸作用动态变化的影响研究; 除了气候因子外, 土壤呼吸作用经验模型应该纳入生物因子等其它影响因素作为变量, 用以提高模型模拟的正确性和准确性。

关键词 土壤呼吸作用 时空变异性 影响机制 生物因子 经验模型

REVIEW OF SPATIAL AND TEMPORAL VARIATIONS OF SOIL RESPIRATION AND DRIVING MECHANISMS

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Abstract Detailed information on the spatial and temporal variation of soil respiration and controlling factors in different terrestrial ecosystems is critical for understanding the ecosystem carbon budget and the response of soils to global climate change. Despite the importance of this topic, knowledge is limited, and key processes and mechanism need clarification. We reviewed recent research advances in the spatial and temporal variation of soil respiration, driving mechanisms and simulation. Environmental and biotic factors play key roles in regulating the temporal variations of soil respiration. Soil respiration also exhibits high levels of spatial heterogeneity, especially across small spatial scales at different time scales. The heterogeneity of vegetation cover, root distribution, major environmental factors and soil properties contributes to the spatial variation of soil respiration. Biotic factors have also been shown to have an effect on soil respiration. However, empirical models of soil respiration typically use soil temperature, soil moisture and their interaction for large-scale soil respiration estimates. Thus, significant errors may result from these models when changes in other biotic factors can confound the temperature or moisture dependence of soil respiration. Therefore, in order to accurately estimate soil respiration in target ecosystems, we must be able to account for its small-scale spatial variation and address the influence of biotic factors in explaining the variation of soil respiration at different temporal scales. Besides climatic variables, it is necessary to incorporate additional factors (biotic factors or soil properties) into

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these empirical models for accurately evaluating soil respiration.

Key words soil respiration, spatial and temporal variations, driving mechanisms, biotic factors, empirical models

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在全球变暖的大背景下,全球碳循环已成为地球科学、生物科学和社会科学共同关注的3个主题之一(陈泮勤,2004)。研究碳循环调控机制并遏止温室气体浓度的持续升高成为国际地圈-生物圈计划(IGBP)、世界气候研究计划(WCRP)、全球环境变化的人文因素计划(IHDP)和生物多样性计划(DIVERSITAS)的核心任务。陆地生态系统碳循环是全球碳循环的重要组成部分,人类活动对陆地生态系统碳循环的影响和干扰日趋严重,这正是当今和未来全球气候变暖的根源(李克让,2002)。

土壤呼吸作用是土壤碳库向大气层碳库输入的主要途径(Schimel, 1995),主要包括来自植物根系的自养呼吸作用和土壤微生物异养呼吸作用。土壤呼吸速率相对微小的改变都会显著改变大气中CO₂的浓度和土壤碳的累积速率,从而加剧或减缓全球气候变暖(Schlesinger & Andrews, 2000)。据预测,未来100年,全球地表温度可能会升高1.6~6.4℃(IPCC, 2007),这会导致土壤呼吸速率在其它因子如土壤湿度、土壤有机质等不变的情况下相应增加(Grace & Rayment, 2000),释放出更多的CO₂,进一步加剧全球气候变暖(Sánchez *et al.*, 2003)。因此,测定不同陆地生态系统土壤呼吸速率及其时空波动,阐明土壤CO₂释放量的环境因子和人为因子的影响,对于全球碳素平衡预算和全球变化潜在效应估计是最为基本的数据(Raich & Schlesinger, 1992)。

然而,土壤呼吸作用又是碳循环最主要环节和认识较为薄弱的部分,有关土壤呼吸作用的影响因素及生态系统间土壤呼吸作用变异性的知识仍存在局限性,一些关键的过程和机制还有待阐明(Maestre & Cortina, 2003; 方精云和王娓,2007)。因此,要准确估算陆地生态系统碳收支,理清气候系统的反馈作用就必须加强土壤呼吸过程及其影响因素的研究(Raich & Tufekcioglu, 2000; Janssens & Pilegaard, 2003),这不仅是准确评估全球碳收支的关键,亦是制定应对全球变化策略的关键。

1 土壤呼吸作用的产生机理

碳以CO₂的形式从土壤向大气圈的流动是土壤呼吸作用的结果(李玉宁等,2002)。土壤呼吸作用,严格意义上讲是指未受扰动的土壤中产生CO₂的所有代谢作用(Singh & Gupta, 1977),包括3个生物学过程(植物根呼吸、土壤微生物呼吸及土壤动物呼吸)和一个非生物学过程(含碳物质的化学氧化作用)。一般认为,土壤呼吸作用主要来自于土壤微生物对有机质(土壤有机质、枯枝落叶和死根等)的分解(异养呼吸作用, RH)及植物根系呼吸(自养呼吸作用, RA)两大部分。有机质进入土壤后,在微生物酶的作用下发生氧化反应,彻底分解而最终释放CO₂、H₂O和能量,即有机质的矿化过程。根系呼吸是指根部及其衍生的呼吸,包括活根组织呼吸、共生的根际真菌和微生物呼吸、根分泌液和死根的分解等活动产生CO₂的过程(Wiant, 1967)。土壤呼吸速率决定了土壤中碳素周转速度,也反映了凋落物碎屑的生产与输入特点和根系呼吸量的水平。

2 土壤呼吸作用的时间变异性及其影响因素

土壤呼吸速率主要受土壤生物区中CO₂产生速率的控制,同时还受影响CO₂在土壤中转移的环境因子的控制(Raich & Schlesinger, 1992)。研究表明,土壤温度、土壤湿度、根系生物量、凋落物、微生物种群、根系氮含量和土壤质地等是影响土壤呼吸作用的重要因素(Boone *et al.*, 1998; Buchmann, 2000; Fang & Moncrieff, 2001; Sánchez *et al.*, 2003; Dilustro *et al.*, 2005)。这些环境因子和生物因子的时间变异性共同驱动着土壤呼吸作用的日变化、季节变化和年际变化,其变化格局强烈地影响着陆地生态系统碳汇强度和大气CO₂浓度。

土壤呼吸作用的日变化多呈单峰型曲线,与土壤温度的变化趋势一致。土壤呼吸作用日变化主要受土壤温度控制,而其它重要的环境变量如

土壤湿度、生物量和土壤性状等在一天内的变化相对较小, 对土壤呼吸作用的影响不明显(Han *et al.*, 2007a)。土壤呼吸速率与土壤温度之间多呈正相关关系(Davidson *et al.*, 2000a; Fang & Moncrieff, 2001; Reth *et al.*, 2004; Jia *et al.*, 2007)。微生物呼吸作用和根系呼吸作用都对土壤温度的变化很敏感(Rey *et al.*, 2002)。在一定范围内环境温度升高可增强微生物活性, 加速土壤中有机质的分解, 从而增加土壤中CO₂浓度; 另一方面, 土壤温度直接影响植物生长和生理活动, 从而影响根系呼吸(Rey *et al.*, 2002); 另外, 土壤温度影响土壤中CO₂向大气的输送过程, 土壤温度升高, CO₂向大气的排放增强(Tang *et al.*, 2003)。

土壤呼吸作用的季节变化主要受气候因子和植被生长控制。从季节看, 夏季土壤呼吸速率高于冬季, 丰水期的土壤呼吸速率高于干早期(王庚辰等, 2004; 任秀娥等, 2007)。夏季植物光合作用强烈, 增进光合产物向土壤的输送, 同时较高的土壤温度亦增强土壤微生物和根系的活性, 从而促进土壤呼吸作用(陈泮勤, 2004); 在热带地区, 土壤温度的季节变化较小, 雨季和旱季的交替决定着土壤呼吸作用的季节变化(Tang & Baldocchi, 2005; Kosugi *et al.*, 2007)。从发育期看, 植物生长旺盛期土壤呼吸速率高于植物发育初期和后期(韩广轩等, 2006)。例如, 作物营养和生殖生长最旺盛时, 土壤呼吸作用也达到峰值(孟磊等, 2005); 羊草群落土壤呼吸的季节变化规律与地上绿色体生物量的季节动态同步(王娓和郭继勋, 2002); 植被群落一般具有明显的季相变化, 季相不同, 对应的群落生物量、碳素和分配同化能力、根系的数量和活性也存在显著差异(Li *et al.*, 1987), 从而加剧土壤呼吸速率的季节变化。

3 土壤呼吸作用空间异质性及其影响因素

土壤呼吸作用在不同时间尺度上具有明显的空间异质性, 特别在森林、草原、农田和荒漠生态系统较小的空间尺度上(Xu & Qi, 2001; Franklin & Mills, 2003; Pen-Mouratov *et al.*, 2006)。Wiseman 和 Seiler (2004)发现人造松林(*Pinus taeda*)中较高的土壤呼吸速率通常出现在靠近树木的地方; 桉树(*Eucalyptus*)人工林中断线附近的土壤呼吸值比行间的土壤呼吸值高(Epron *et al.*,

2004); 靠近松树的土壤呼吸速率显著高于远离松树的土壤呼吸速率(Pangle & Seiler, 2002); Fang等(1998)发现美洲沼泽松(*Pinus caribaea*)下的土壤呼吸速率显著高于空旷地上的土壤呼吸速率; 较高的土壤呼吸速率通常出现在靠近玉米(*Zea mays*)植株的地方, 土壤呼吸速率靠近植株>株间>行间(Han *et al.*, 2007b)。

植被覆盖、根系分布、主要的环境因素和土壤特性的异质性导致了土壤呼吸作用的空间异质性, 因为土壤水分、根系生物量、微生物数量或者凋落物量随着距离植株的远近而发生变化(Stoyan *et al.*, 2000; Franklin & Mills, 2003; Epron *et al.*, 2004; Wiseman & Seiler, 2004)。Stoyan等(2000)利用地统计学方法分析了2 m²尺度上白杨(*Populus*)林和小麦(*Triticum aestivum*)地中土壤呼吸作用和土壤性质的空间异质性, 认为土壤呼吸作用的空间异质性部分是由于植株根系和植株凋落物的空间格局造成的。在典型人工林中, 根系生物量通常呈现一定的辐射梯度(Wiseman & Seiler, 2004)。在这个辐射梯度中, 靠近植株的根系生物量大于远离植株的根系生物量, 因此距离植株越远, 根系呼吸值越小(Han *et al.*, 2007b)。因此, 土壤呼吸作用的空间异质性可能与根系密度有关(Shibistova *et al.*, 2002)。

另外, 一个区域内部或者不同区域间的土壤呼吸作用的空间异质性也可能由土壤湿度、土壤质地和土壤化学性质的差异性引起(Tang & Baldocchi, 2005)。Maestre和Cortina (2003)认为小尺度上的土壤湿度和土壤温度的差异性导致了土壤呼吸作用的空间异质性。Stoyan等(2000)发现在人工白杨(*Populus*)林中, 由于雨水顺着树干流下, 使得树干周围的土壤湿度较大, 这是导致树干周围的土壤呼吸速率较高的原因。松树(*Pinus ponderosa*)林中土壤总氮、磷、镁和有机质的空间异质性可以解释土壤呼吸作用空间异质性的44%~55% (Xu & Qi, 2001)。

4 土壤呼吸作用模拟研究

目前, 对于如何模拟环境因子对土壤呼吸作用的影响仍然存在着激烈的争论(Buchmann, 2000; Pumpanen *et al.*, 2003)。大多数研究者常应用经验模型模拟土壤呼吸作用强度及其时间变化, 由于土壤环境的复杂性, 过程模型很少应用

(Michelsen *et al.*, 2004)。这些经验模型通常使用土壤温度(Fang *et al.*, 1998; Buchmann, 2000; Janssens & Pilegaard, 2003)、土壤湿度(Davidson *et al.*, 2000b; Epron *et al.*, 2004; Sotta *et al.*, 2004)或者两者的交互作用(Tufekcioglu *et al.*, 2001; Lee *et al.*, 2002; Tang & Baldocchi, 2005)估算大尺度的土壤呼吸作用。例如,研究者建立了多种方程用以描述土壤呼吸作用与土壤温度的关系,包括线性方程(Fan *et al.*, 1995; O'Connell *et al.*, 2003; Chimner, 2004)、指数方程(Buchmann, 2000; Sánchez *et al.*, 2003; Reth *et al.*, 2004; 杨继松等, 2008)、Arrhenius方程(Lloyd & Taylor, 1994; Thierron & Laudelout, 1996)、幂函数方程(Fang & Moncrieff, 2001)和逻辑斯缔方程(Jenkinson, 1990; Rodeghiero & Cescatti, 2005)。目前,许多生态系统模型用指数函数描述土壤呼吸速率和温度的关系(Langley *et al.*, 2005),土壤呼吸作用的温度敏感性(Q_{10})成为大气碳平衡估算中的一个关键参数(Raich & Schlesinger, 1992)。越来越多的证据表明, Q_{10} 不是一个常数,而随着土壤湿度、根系生物量、凋落物和微生物数量等环境因子的季节变化而变化(Davidson *et al.*, 1998)。另外,土壤水分直接参与生物的生理过程,其在过低或过高的水平下会限制土壤呼吸作用(Pangle & Seiler, 2002)。土壤呼吸速率与土壤水分含量之间的关系也可用多种方程函数来描述,常用的有线性方程(Davidson *et al.*, 1998)、二次方程(Mielnick & William, 2000)、指数模型(Keith *et al.*, 1997)、对数模型(Davidson *et al.*, 2000b)和双曲线方程(Schlentner & Cleve, 1985)。

然而,土壤呼吸作用不单纯是对土壤温度的生理响应过程,而是几个复杂生态系统过程共同作用的结果(Janssens & Pilegaard, 2003)。当这种复合关系存在时,不可能单独分离出土壤呼吸作用的温度或湿度效应(Davidson *et al.*, 1998; Janssens & Pilegaard, 2003)。例如,秋天随着土壤温度下降,土壤呼吸速率反而会继续增大,这是因为秋季落叶和细根的分解可以提供土壤微生物更多的基质(Davidson *et al.*, 1998)。除了土壤温度和湿度,光合速率、净初级生产力(NPP)、根系生物量、凋落物和微生物种群等生物学特性也是影响土壤呼吸作用的重要因素(Boone *et al.*, 1998; Buchmann, 2000; Fang & Moncrieff, 2001;

Sánchez *et al.*, 2003; Dilustro *et al.*, 2005)。

生物因子通过影响土壤微气象和结构、凋落物数量和质量、根系呼吸作用,进而影响土壤呼吸作用(Raich & Tufekcioglu, 2000)。植物生长期问,土壤呼吸作用主要受植物生长控制(Högberg *et al.*, 2002; 杨兰芳和蔡祖聪, 2004),主要表现在:1) 土壤呼吸作用的物质基础来源于植物光合作用(Rochette & Flanagan, 1997; Lohila *et al.*, 2003),光合作用强烈时,地下的呼吸作用也旺盛(Atkin *et al.*, 2000)。根系呼吸作用主要依赖于植物地上部分光合产物对地下部分的分配(Yuste *et al.*, 2004),分配到根系中的光合产物约75%被呼吸消耗掉,只有25%用于树木生长(Högberg *et al.*, 2002);在玉米(*Zea mays*)生长期问,土壤呼吸主要来自于新近合成的光合产物(杨兰芳等, 2007);Högberg等(2001)证实土壤呼吸速率与植被光合作用呈正相关;大麦(*Hordeum vulgare*)田中,土壤呼吸速率与植被总第一性生产力(GPP)呈显著正相关(Moyano *et al.*, 2007)。2) 植物根系是土壤呼吸作用的主要参与者,其根量与根系活性决定土壤呼吸作用的强弱。在不同陆地生态系统中,根系呼吸作用占土壤呼吸作用的比例大部分在10%~90%之间(Hanson *et al.*, 2000),这些比值依赖于植物生长阶段及气候条件等。Ben-Asher等(1994)研究发现土壤呼吸与根系的主要特性之间具有满意的的相关性。根系的参与极大地促进土壤呼吸,根系呼吸速率随生物量的增加而增加(孙文娟等, 2004)。3) 生物量通过影响土壤中凋落物和碎屑的数量来影响微生物的生长和活性,从而影响土壤呼吸作用(Kuzyakov & Cheng, 2001; Lohila *et al.*, 2003)。研究发现,加拿大北部森林所有采伐迹地地下碳通量与立地现存生物量呈正相关(Pypker & Fredeen, 2003);近熟林土壤呼吸随着凋落物的增加而增加(Raich & Nadelhoffer, 1989)。4) 在生态系统尺度上,植物NPP是控制土壤生物和地下过程的最重要因子(Wardle, 2002)。土壤呼吸作用与NPP在植被尺度上存在着线性关系(Raich & Schlesinger, 1992),有充分的证据表明植物生长速率与土壤呼吸作用是紧密联系的过程(Raich & Tufekcioglu, 2000)。NPP决定着植物地上和地下凋落物碎屑向土壤中输入碳量的水平和土壤微生物的活性(Raich & Potter, 1995),地下凋落物和碎屑的任何变化都可能强

烈地影响土壤微生物呼吸(Rey *et al.*, 2002)。因而,微生物呼吸作用对生态系统生产力有间接的依赖性。

5 问题与展望

综上所述,由于土壤呼吸作用在碳循环中的重要作用,人们在土壤呼吸作用时空动态变化规律、影响机制和定量评估等方面开展了大量的研究工作,并取得了较大进展。但是目前关于土壤呼吸作用的控制机理及其过程认识还不统一,对于土壤呼吸作用的定量评估仍具有较大的不确定性。具体体现在以下方面:

1) 土壤呼吸作用空间异质性及其影响机制研究不足。陆地生态系统中土壤呼吸作用主要有3个方面的特征:土壤呼吸强度、时间变异性与空间异质性(Fang *et al.*, 1998)。目前,主要集中于土壤呼吸作用强度、时间变异性及其影响因素影响方面,关于土壤呼吸作用的空间异质性及其影响因子的研究较少。要精确估算生态系统的碳收支,必须解决土壤呼吸作用小尺度上的空间变异性(Xu & Qi, 2001; Adachi *et al.*, 2004),如果不考虑土壤呼吸作用的空间异质性就把田间测定的土壤呼吸作用尺度化到生态系统层面,将会导致很大的偏差(Tang & Baldocchi, 2005)。

目前,土壤呼吸作用测定中,动态气室-红外CO₂分析仪法(IRGA)使用最为广泛,如Li-cor系列(6250, 6400, 8100等)。在测量过程中,为了避免对土壤的扰动,一般把气室放置在土环上。以前的土壤呼吸作用测定中,土环的位置多描述为任意放置(Buchmann, 2000; Sánchez *et al.*, 2003)。但是在人工林和农田生态系统中,由于植株空间布局的规律性,应该考虑植株的空间格局对土壤呼吸作用的影响(Han *et al.*, 2007b)。因此,为了精确估算土壤呼吸作用,应该考虑生物因子和土壤特性的空间异质性,但是目前定量评估土壤呼吸作用的空间异质性仍然有限和困难(Rayment & Jarvis 2000; Tang & Baldocchi, 2005),这是未来土壤呼吸作用研究中的一个关键且具有挑战的领域(Maestre & Cortina, 2003)。

2) 生物因子对土壤呼吸作用的影响机制尚不清楚。土壤呼吸作用是几个不同过程共同作用的结果,包括复杂的生物过程,任何单一过程的变化都会掩盖其它过程的作用(Buyanovsky &

Wagner, 1995)。尽管研究者已经认识到生物因子强烈地影响着土壤呼吸作用的强度和动态变化(Raich & Tufekcioglu, 2000; Högberg *et al.*, 2001, 2002; Wardle, 2002),但是关于生物要素对土壤呼吸作用的影响机制还缺乏深入理解。人们习惯通过对气候因子的测定来对土壤呼吸作用进行定量评估,但是当生物因子发生变化影响土壤呼吸作用的温度或水分效应时,仅依靠气候因子模拟土壤呼吸作用,缺乏生理学基础,具有明显的局限性(Subke *et al.*, 2003),会导致明显的偏差和错误(Janssens & Pilegaard, 2003)。因此,应该对不同区域典型生态系统进行长期定位观测,加强土壤呼吸作用与生物量、光合作用和净初级生产力等生物要素的同步观测,增进生物因子对土壤呼吸作用的影响过程与机理的理解。

3) 耦合生物因子和水热因子综合影响的土壤呼吸作用动态模型还鲜有报道。研究者已经建立了土壤呼吸作用数据库并发展了模型,用以把田间测定的土壤呼吸作用尺度化到生态系统或者更大的空间尺度。这些模型通常利用土壤温度、土壤湿度或者两者的交互作用模拟土壤呼吸作用。但是,这些经验模型很难揭示土壤呼吸作用与控制其时空变化之间的内在规律(方精云和王娓, 2007)。除了气候因子外,这些经验模型应该纳入生物因子等其它影响因素作为变量,用以提高模型模拟的正确性和准确性(Högberg *et al.*, 2001; Han *et al.*, 2007a, b)。另外,尽管土壤温度和湿度可以很好地模拟土壤呼吸作用的时间变异性,但是它们并不能充分解释土壤呼吸作用在样地内部或者样地间的空间异质性(Xu & Qi, 2001; Tang & Baldocchi, 2005)。因此,为了把箱式法测定的土壤呼吸作用尺度化到生态系统层面,建立一个既能反映土壤呼吸作用时间变异性又能反映土壤呼吸作用空间异质性的模型是十分必要和紧迫的(Xu & Qi, 2001; Han *et al.*, 2007b)。

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